

Paleomagnetic Pole Positions and Geomagnetic Secular Variation from the Cretaceous Ponta Grossa Dike Swarm (Brazil)

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Received: February 19, 2014; accepted: July 03, 2014; published on line: March 31, 2015

Resumen

A detailed paleomagnetic and rock magnetic investigation has been carried out on the Early Cretaceous Ponta Grossa dike swarm in southern Brazil. This formation seems an excellent target for paleomagnetic study. The dikes are widely distributed over a large area, easy to access, and they record faithfully the geomagnetic field at the time of the eruption. Most of them are fresh and have been dated by K-Ar and Ar-Ar. Thermomagnetic experiments (low-field versus temperature curves) suggest low-Ti titanomagnetites as main remanence carriers, and their domain structure is characterised by a mixture of single-domain and multi-domain grains. Characteristic paleomagnetic directions are retrieved from 28 out of 29 sites (235 standard paleomagnetic cores). 17 sites show normal polarity, 10 sites show reverse polarity and one site shows an oblique direction, with negative inclination, separated 79° of

the mean normal directions. The reversal test is positive at the 95% confidence level which ensures that the secondary remanent magnetizations were successfully removed and the sampling adequately averaged the palaeosecular variation. The paleosecular variation parameters values obtained in this study correlate with those determined for the Cretaceous Normal Superchron, between 125 to 84 Ma. The new paleomagnetic pole position (222.0° E, 88.1° S, $N = 24$, $K = 35.68$, $A_{95} = 5.0^\circ$) agrees well (within uncertainties) with the reference poles determined from Besse and Courtillot (2002), and disagree with those reported in previous studies. The new results should be considered for estimating the Cretaceous paleomagnetic poles for stable South America.

Palabras clave: Ponta Grossa Dike Swarm, South America, Paleomagnetism, Cretaceous.

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Abstract

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Key Words: Ponta Grossa Dike Swarm, South America, Paleomagnetism, Cretaceous.

Introduction

The study of the remanent magnetism recorded in rocks is the basis for constraining important problems in plate tectonics and paleogeographic reconstructions at different time scales. Paleomagnetic poles are often used to determine past location of continents. Reliable reconstructions of geological terrains cannot be obtained from poorly defined paleomagnetic poles, no matter how many sites are used. Precisely defined reference paleopoles (Besse and Courtillot, 2002) are crucial to produce robust paleocontinental reconstructions.

The Parana Magmatic Province and surrounding volcanic areas, including the Ponta Grossa Dikes have been subject of many paleomagnetic studies (Ernesto *et al.*, 1990, 1996, 1999; Raposo *et al.*, 1995; Raposo *et al.*, 1998; Alva-Valdivia *et al.*, 2003; Mena *et al.*, 2006). However, the paleomagnetic pole positions obtained in these studies present unusually high dispersion, which probably do not reflect the true characteristics of the Earth's magnetic field for the time interval of ~140–120 Ma. As discussed by Alva-Valdivia *et al.* (2003), the apparent high scatter in paleomagnetic pole positions may be due to unrecognized tectonic disturbances (see also Ernesto *et al.*, 1990).

On other hand, the fine scale characteristics of the geomagnetic field during the Cretaceous is of great interest in understanding the long-

term evolution of the geomagnetic field, related to the constant polarity superchron, the polarity reversal frequency patterns and paleosecular variation. The fluctuation of earth's magnetic field elements is closely connected to conditions in the Earth's liquid core and at the core-mantle boundary. Paleomagnetic data provide an invaluable source of information on the Earth's liquid core and deep interior.

There are two clear intervals in the geological past when the Earth's magnetic field remained with constant polarity for a considerable length of time: from about 125 to 84 Ma the polarity was normal (so-called the Cretaceous Normal Superchron—CNS) and from about 310 to 260 Ma the polarity was reversed (so-called the Permo-Carboniferous Superchron—PCRS). During the last decades, many investigations were devoted to study the morphology of the earth's magnetic field during the Cretaceous (Biggin *et al.*, 2008). Namely, most important question is related to the variability during and around the CNS. The particularity of this interval (125 to 84 Ma after Cande and Kent, 1995 and Tarduno *et al.*, 2002) is that Earth's magnetic field did not reverse polarity. Knowledge of the geomagnetic field behaviour just prior to CNS is patchy and largely unknown. Thus, more accurate paleomagnetic studies giving information about the past behavior of the geomagnetic field are strongly needed.

High quality paleomagnetic reference poles are required to distinguish true polar wander effects from relative hotspot motion, and to

constrain major geodynamic events (Torsvik *et al.*, 2002; Somoza and Zaffarana, 2008). Major tectonic events characterize the mid- and late Cretaceous, with the final stages of Gondwana breakup, volcanic plateau emplacement and fast seafloor spreading rates (e.g., Taylor, 2006; Seton *et al.*, 2009).

The present study is aimed to determine precise Early Cretaceous paleomagnetic pole positions for stable South America, and contribute to the study of the fluctuations of the Earth's magnetic field just prior to the Cretaceous Normal Superchron. A preliminary paleointensity study was already reported from the Ponta Grossa dikes, analyzing 3 sites with 11 samples (Cejudo *et al.*, 2009). The present study expands on the initial reconnaissance study, based on 335 standard paleomagnetic cores belonging to 29 dikes.

Geological Setting and Sampling

The Parana-Etendeka area represents one of the largest known continental flood basaltic provinces. It is constituted of volcanic and sub-volcanic rocks, with the majority being tholeiitic basalts and andesites with subordinate rhyolites, covering an area of around $1.2 \times$

10^6 km² in southern Brazil (mainly), northern Argentina, Uruguay and Paraguay. Based on available radiometric data, it appears that these huge magmatic eruptions occurred between 133 and 132 Ma (Renne *et al.*, 1992, 1996).

Intra-continental volcanism that crop outs around the basaltic traps at the present time is younger, and comprises basaltic and andesitic dikes. The most important mafic dike swarms in Brazil occur in the Ponta Grossa (PG) region (Figure 1) and are associated with the flood basalt suites of the Paraná basin (Piccirillo *et al.*, 1990; Brandt *et al.*, 2010). The region is cut by hundred of dikes, predominantly basaltic and andesitic composition. There is now a general agreement that PG dikes were probably feeders of the stratovolcanoes built in northern Paraná (Sial *et al.*, 1987). Renne *et al.* (1996) reported a comprehensive geochronological study provided numerous ⁴⁰Ar/ ³⁹Ar high quality plateau ages. The age-probability distribution for the dominant pulse (131.4 ± 0.4 to 129.2 ± 0.4 Ma) shows a pronounced peak at 130.5 Ma. Thus, Ponta Grossa dikes are younger than the southern Paraná Magmatic Province lavas, which occurred at 133-132 Ma (Renne *et al.*, 1996).

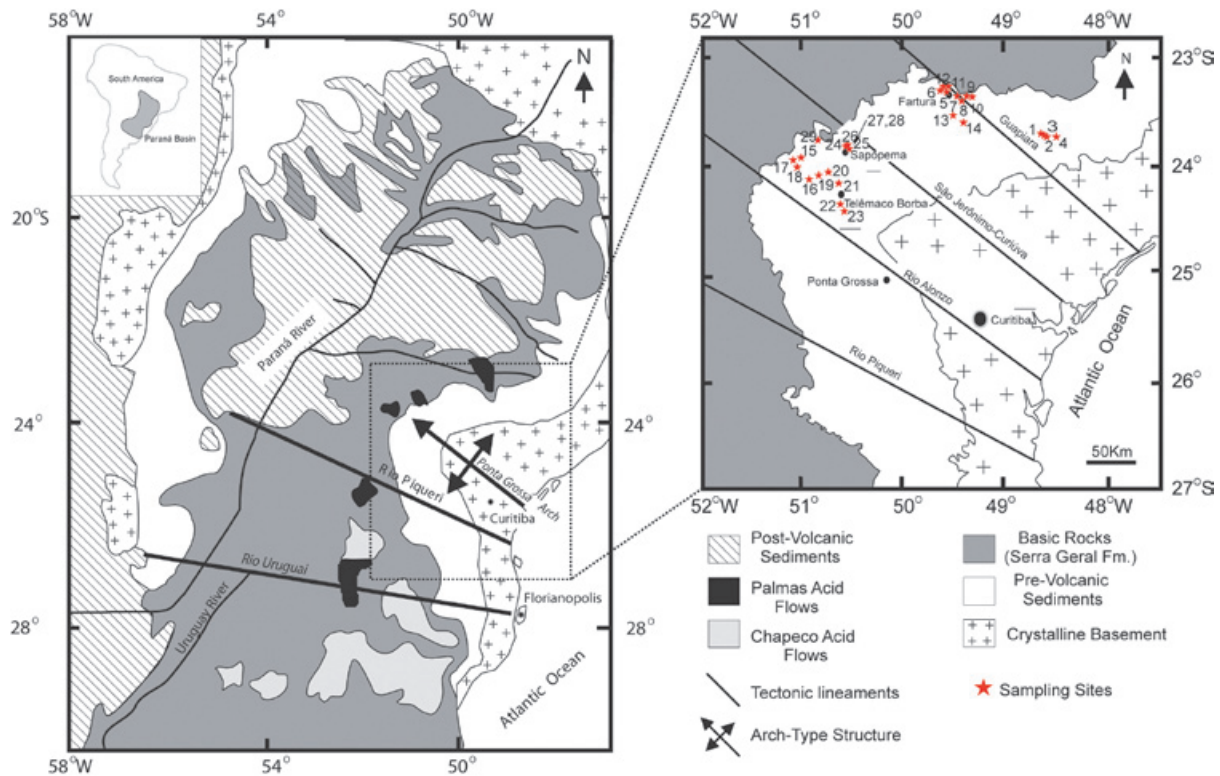


Figure 1. Simplified geologic map of Southern Brazil showing the location of studied sites (adopted from Brandt *et al.*, 2009 and Cejudo *et al.*, 2009).

We obtained 235 standard paleomagnetic cores belonging to 29 sites distributed along road outcrops in Ponta Grossa region, Southern Brazil (Figure 1). The samples were distributed throughout each dike both horizontally and vertically. In general, samples were obtained at least 30 cm distance from the dike edge (when possible). Cores were obtained with a gasoline-powered portable drill, and then oriented in most cases with both magnetic and sun compasses.

Laboratory Procedures

Magnetic experiments were carried out to find out the carriers of remanent magnetization, to estimate their domains states and grain sizes and to obtain information about their paleomagnetic stability. From each of the 29 studied dikes, one sample was selected for these experiments, which included the measurement of low-field susceptibility versus temperature (k - T) curves, the determination of hysteresis parameters and the measurement of isothermal remanent magnetization (IRM) acquisition curves. All hysteresis measurements were carried out with the AGFM-Micromag apparatus, while continuous thermomagnetic curves were obtained using Bartington susceptibility bridge equipped with furnace.

Thermomagnetic curves were measured in air, heating the samples up to 600°C and cooling them down to room temperature. Curie

points (T_C) were determined using the Prevot (1983)'s method. The most representative behaviour is reported on Figure 2a. It displays irreversible curves with a single ferromagnetic phase with a high Curie temperature near 550°C, which corresponds to low-Ti titanomagnetite or slightly Al- or Mg-substituted magnetite. We considered a curve to be reversible if the difference between initial magnetization before heating had started and final magnetization after cooling had been completed was less than $\pm 10\%$.

Hysteresis parameters J_s (saturation magnetization), J_r (saturation remanence), H_c (coercivity) and H_{cr} (coercivity of remanence) were obtained from hysteresis (Figure 2b) and backfield curves. Analysis of the measurements was performed with the RockMagAnalyzer 1.0 software (Leonhardt, 2006). The obtained values point that studied samples can be match to the PSD (pseudo-single-domain) magnetic grains. This behavior might be also explained by a mixture of single-domain (SD) and multi-domain (MD) particles (Dunlop, 2002). IRM acquisition curves (Figure 2b) show that more than 95% of saturation magnetization (SIRM) was reached with applied fields of less than 150 mT. Thus IRM measurements point to low-coercivity phases as main remanence carriers.

Remanent magnetization measurements were performed at the 'Laboratorio Interinstitucional de Magnetismo Natural' in Campus

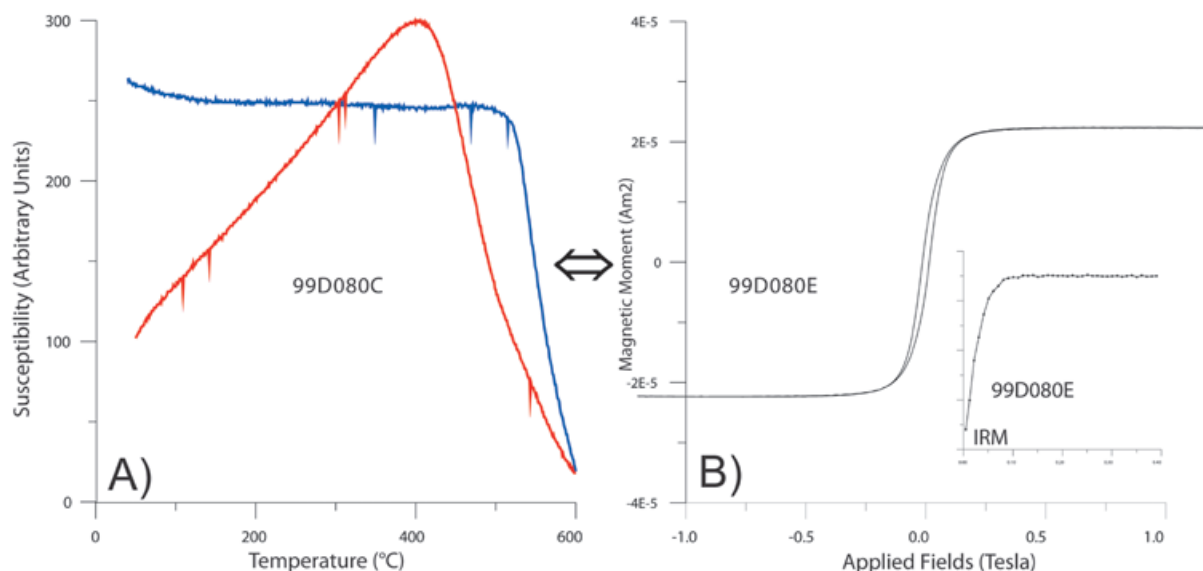


Figure 2. a) Susceptibility versus temperature (in air) curve of representative samples. b) Typical examples of hysteresis loop (corrected for dia/paramagnetism) of small chip samples from the studied volcanic units and associated IRM (isothermal remanent magnetization) curve.

Morelia (UNAM) with a spinner JR6 (nominal sensitivity $\sim 10^{-9}$ Am²) magnetometer. Most of samples were cleaned by means of alternating field demagnetization using a LDA-3 demagnetizer, with 95 mT as maximum fields available. Eleven sites displayed a single paleomagnetic component (Figure 3, samples 99D015A), whereas in the remaining ones a viscous present-day normal-polarity overprint could also be observed that was easily removed (sample 99D010B). Some dikes however yielded relatively strong secondary components (sample 99D115A), which were removed after applying around 60 mT alternating fields. The origin of this overprint is probably related to hydrothermal alteration.

The direction of the characteristic magnetization (ChRM) was determined by the least squares method (Kirschvink, 1980), with 4 to 9 points being taken for the principal component analysis. The directions obtained were averaged by volcanic unit and the statistical parameters calculated assuming a Fisherian distribution.

Main Results and Discussion

17 sites show normal polarity remanences, 10 sites show reverse polarities and one site have an outlier oblique direction, with negative inclination, separated 79° of the mean normal direction (calculated before exclude it; Figure 4). The mean paleomagnetic direction for the

Table 1. Paleomagnetic results of studied samples. *D, I* : declination and inclination of the site mean directions; *N* is number of specimens used; *k* and *a95* are Fisherian statistical parameters; VGP Long, VGP Lat : longitude and latitude of the virtual geomagnetic poles.

Site	Lat (south)	Long (west)	N	D	I	k	a95	Polarity	VGP Long [°E]	VGP Lat [°S]
PG02	24°11.423	48°32.856	7	187,4	42,6	72	7,2	R	227,4	-83,2
PG03	24°11.325	48°32.235	7	167,6	52,5	149	5,1	R	359,3	-76,0
PG04	24°11.269	48°32.654	7	202,8	22,3	162	5,6	R	195,8	-65,0
PG05	23°23.292	49°33.018	8	4,8	-38,9	748	2,8	N	204,6	-85,4
PG06	23°23.524	49°32.753	4	182,3	40,8	47	11,6	R	219,6	-87,9
PG07	23°23.356	49°32.468	5	184,8	19,9	191	5,6	R	150,5	-76,1
PG08	23°23.825	49°32.112	8	191,8	22,3	342	3,3	R	176,0	-73,7
PG09	23°20.072	49°28.498	7	19,7	-50,6	68	7,4	N	249,4	-70,8
PG10	23°21.226	49°28.786	8	339,7	-45,9	79	6,8	N	24,2	-71,3
PG11A	23°22.094	49°29.099	4	226,1	-50,5	58	9,6	O	174,4	-19,8
PG11B	23°22.094	49°29.099	4	15,4	-27,3	299	5,3	N	191,8	-73,0
PG12	23°22.511	49°29.638	6	179,2	30,2	125	6,1	R	124,4	-82,8
PG13	23°26.102	49°26.026	7	352,4	-43,8	258	4,2	N	21,5	-82,8
PG14	23°29.195	49°19.562	6	349,8	-41,8	192	4,8	N	34,9	-80,6
PG15	24°04.002	51°03.810	5	165,2	50,6	61	9,9	R	6,6	-75,0
PG16	24°11.103	51°04.012	5	169,2	43,9	106	7,5	R	27,9	-80,1
PG17	23°59.824	51°05.091	6	178,2	47,2	39	11,5	R	328,9	-85,3
PG18	24°08.105	51°00.181	7	342,9	-45,4	48	9,3	N	25,3	-74,3
PG19	24°09.324	50°58.474	6	12,9	-48,1	69	8,2	N	245,2	-77,5
PG20	24°13.512	50°55.681	6	18,6	-49,9	63	8,5	N	244,6	-72,3
PG21	24°15.715	50°53.593	5	14,6	-50,9	59	7,9	N	252,0	-75,2
PG22	24°26.463	50°44.253	6	11,4	-30,9	47	9,2	N	185,3	-76,8
PG23	24°25.091	50°42.341	5	8,2	-47,7	143	5,6	N	251,9	-81,5
PG24	23°50.687	50°37.158	6	4,6	-32,1	84	7,4	N	134,0	-82,3
PG25	23°50.977	50°37.057	6	359,6	-34,7	66	8,5	N	124,8	-85,2
PG27	23°51.950	50°36.975	5	356,2	-35,4	38	10,8	N	89,3	-84,4
PG28	23°52.474	50°36.601	6	1,7	-41,8	43	9,2	N	227,5	-88,4
PG29	23°54.231	50°35.388	1	344,5	-47,1	N.D	N.D	N	18,5	-75,4

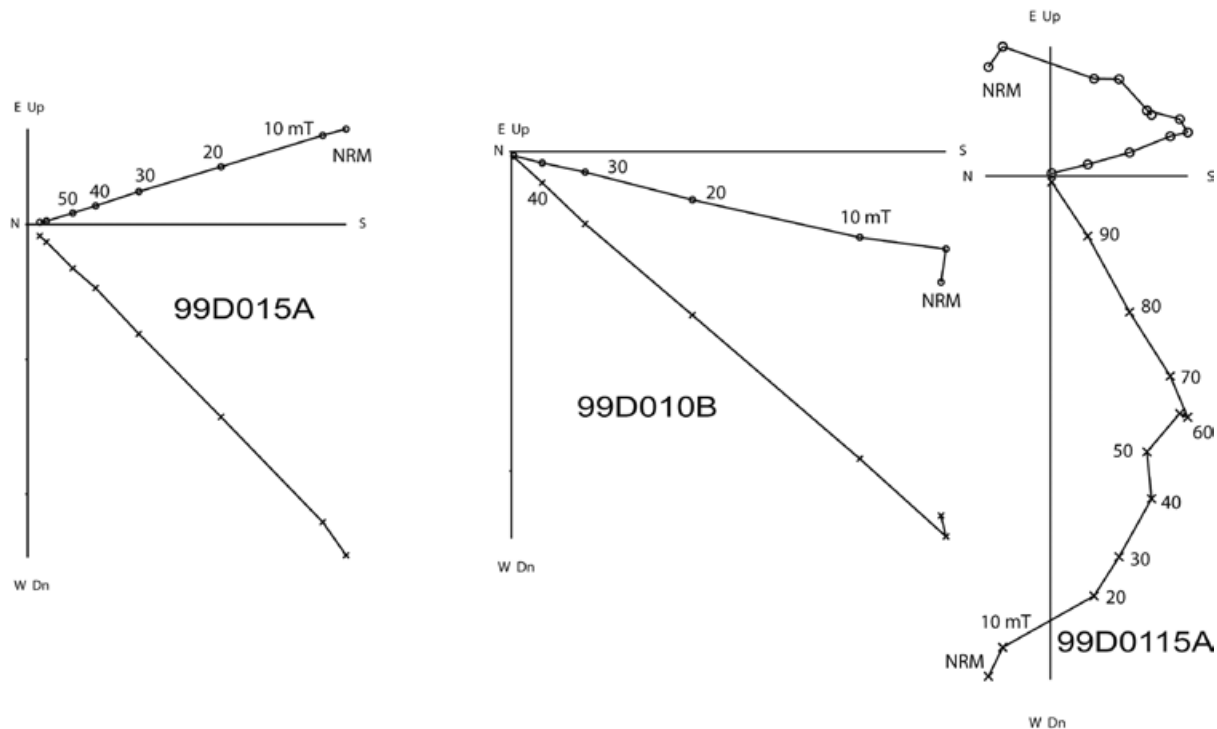


Figure 3. Orthogonal vector plots of stepwise alternative field demagnetization of three representative specimens.

normal polarity sites is $I = -42.6^\circ$, $D = 2.2^\circ$, $k = 46.9$, $a_{95} = 5.3^\circ$, $N = 17$ (k is precision parameter and N number of sites), while reversely magnetized sites give a mean direction with $I = 37.8^\circ$, $D = 182.2^\circ$, $k = 27.62$, $a_{95} = 9.3^\circ$, $N = 10$. The reversal test as defined by McFadden and Lowes (1981) is accepted at the 95% confidence level and the reversal test from McFadden and McElhinny (1990) is positive with classification B ($gc = 9.58^\circ$; $go = 4.8^\circ$). This ensures that the secondary remanent magnetizations were successfully removed and the sampling adequately averaged the palaeosecular variation.

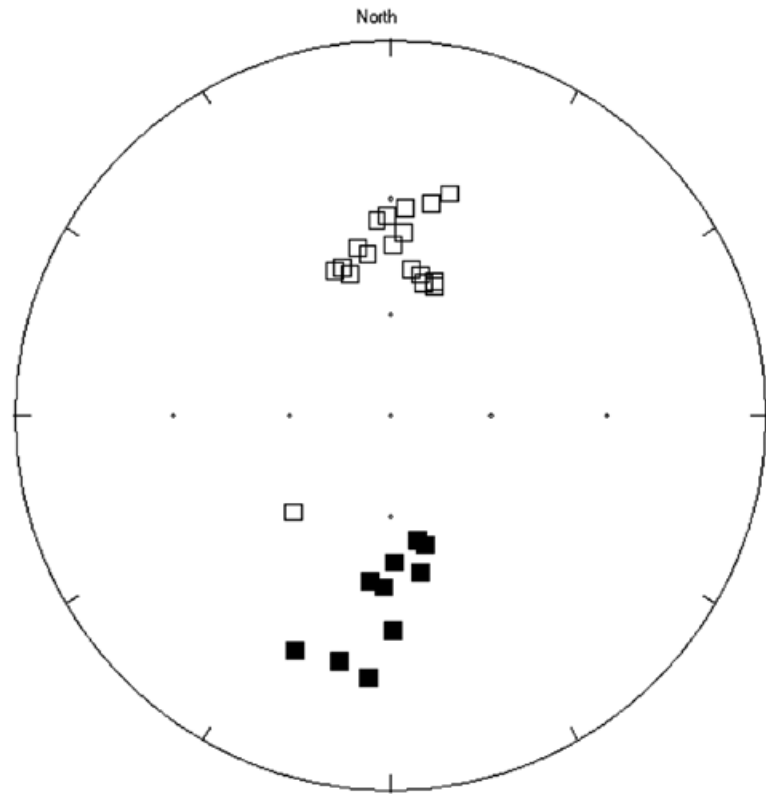
For both sets, the quantil-quantil (Q-Q) plot for declinations using the uniform model is approximately linear, passing through the origin with a slope near 45° (Figure 5). The Q-Q Co-Inclination plot with exponential model is also linear, with a slope that gives an estimated $k = 58.82$. The Q-Q plot for two variables is approximately linear, passing through the origin and with a slope that give an estimate $k = 42.16$. The three k estimate values are statistically undistinguishable at 95% level using an F-test. These graphical tests suggest that the underlying distributions are fisherian.

The distribution of the site-mean VGPs is somewhat elongated (Figure 6), with

maximum elongation direction along the 24° - 204° meridian. The eigenvalues of the VGP orientation matrix are $t_1 = 25.714$; $t_2 = 0.987$; $t_3 = 0.299$, then the elongation index is 3.3. Similar configuration was also found for coeval VGPs (Mena *et al.*, 2006, 2011). However, we consider the elongation not significant, since the VGP distribution yields a reasonably good fit with fisherian distribution when probability plots as well as formal testing procedures are used.

The VGPs quantil-quantil (Q-Q) plots are consistent with a fisherian distribution (Figure 7), with slopes that give estimated $k = 43.5$ from colatitudes and $k = 73.1$ from two variables. Using an F-test these values are statistically undistinguishable, at 95% confidence level, of the value $k = 39.6$ estimate with the VGPs. Employing formal testing procedures, we obtain the following results for the modified statistics of Kolmogorov-Smirnov and Kuiper (Dn^* , Vn^* , $MU(Vn)$, $ME(Dn)$ and $MN(Dn)$; Fisher *et al.*, 1987): Longitude test: $Dn^* = 0.8677$; $Vn^* = 1.2862$; $MU(Vn) = 1.1780$; colatitude test: $Dn^* = 0.6381$; $Vn^* = 1.0538$; $ME(Dn) = 0.6162$; Two-variable test: $Dn^* = 0.7390$; $Vn^* = 1.3542$; $MN(Dn) = 0.7408$. All values suggesting no departure from the Fisher model.

Figure 4. Stereoplot for site mean paleodirections corresponding to all 28 sites where reliable results are obtained.



Considering the PGVs from sites with $\alpha_{95} < 10^\circ$ (excluding PG6, PG17 and PG27 and the oblique site PG11a) the paleomagnetic pole (PP) is located at 222.0° longitude E, 88.1° latitude S, $N = 24$, $K = 35.68$, $A_{95} = 5.0^\circ$. The

inclusion of all sites into the mean calculation gives: 222.4° longitude E, 88.4° latitude S, $N = 27$, $K = 39.61$, $A_{95} = 4.5^\circ$, which is not significantly different.

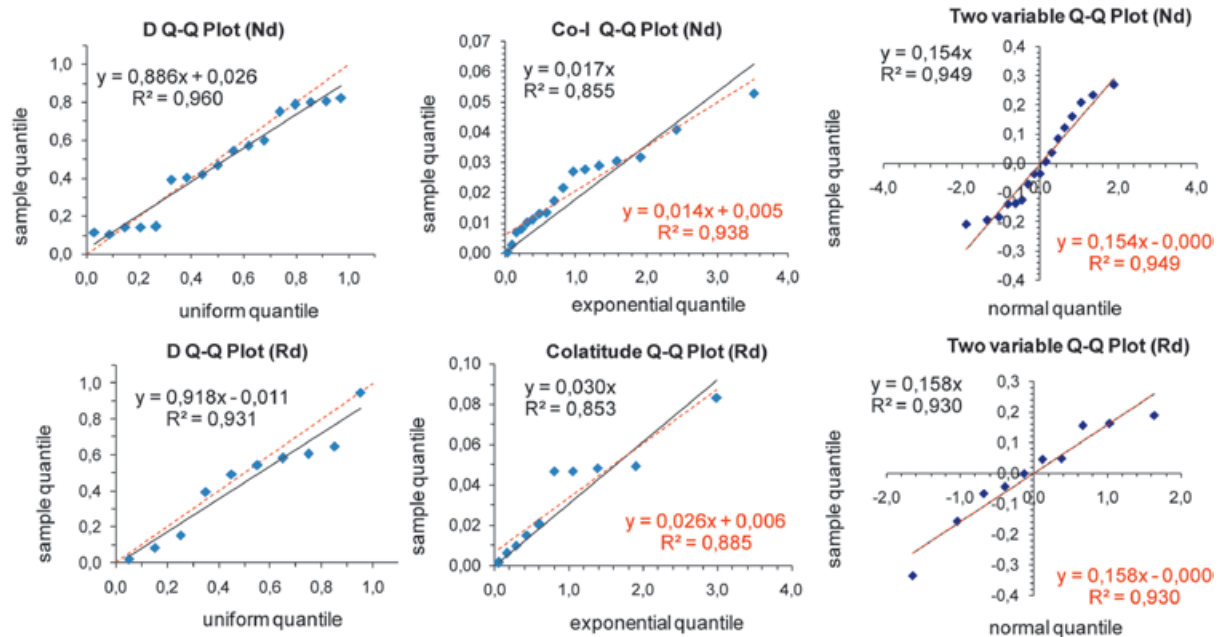


Figure 5. Probability plots for checking goodness of fit of the paleodirections to the Fisher distribution.

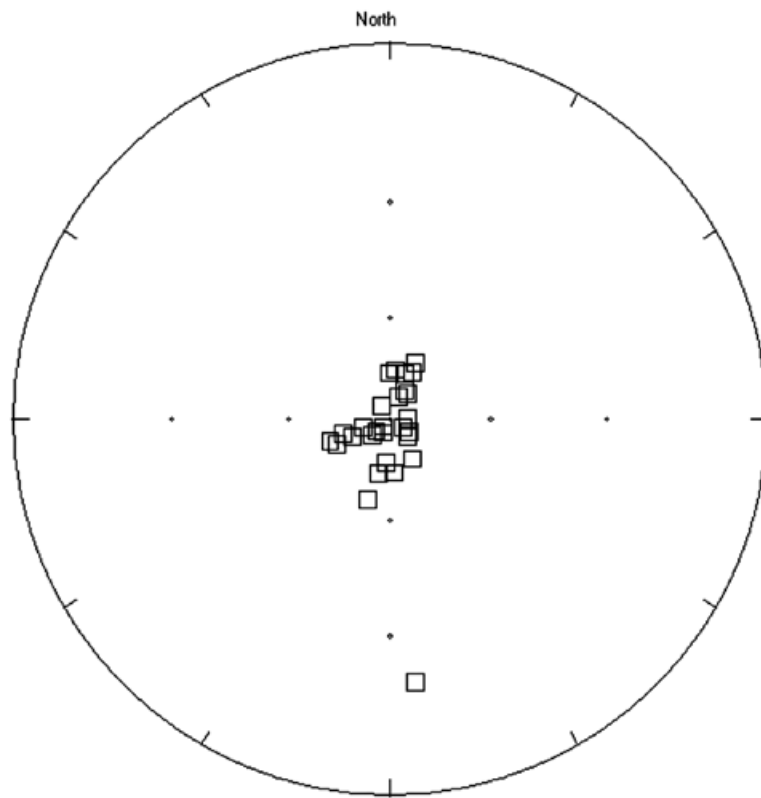


Figure 6. South VGPs distribution from 28 sites.

The mean paleomagnetic pole obtained in this study is shown in the present geographical reference frame in Figure 8 and listed in Table 2 together with previously published South American poles. In general, the pole obtained in this study, agrees reasonably well with other pole positions, in particular with Misiones (Argentina), Central Paraná Magmatic Province, Florianópolis Dikes (Brazil), Los Adobos (Geuna *et al.*, 2000) poles. However, some other similar age paleomagnetic poles (Central and Northern Paraná Basin in Brazil and Arapey Formation in Uruguay) are clear outliers. As initially suggested by Ernesto *et al.* (1996) and recently confirmed by

Goguitchaichvili *et al.* (2013), this may be due to some local tectonic rotations or insufficient sampling to average the paleosecular variation as well as to unrecognized structural disturbances. Overall, the Ponta Grossa poles are significantly different from the pole position estimated from the hotspot reconstruction (Muller *et al.*, 1993). This fact is attributed by several authors (i.e. Alva-Valdivia *et al.*, 2003, Cervantes *et al.*, 2010, Goguitchaichvili *et al.*, 2013) as cause of the true polar wander or hotspot motion. Somoza and Zaffarana (2008) analyzed the mid-Cretaceous poles for South America, concluding that there is no detectable motion with respect of the paleomagnetic

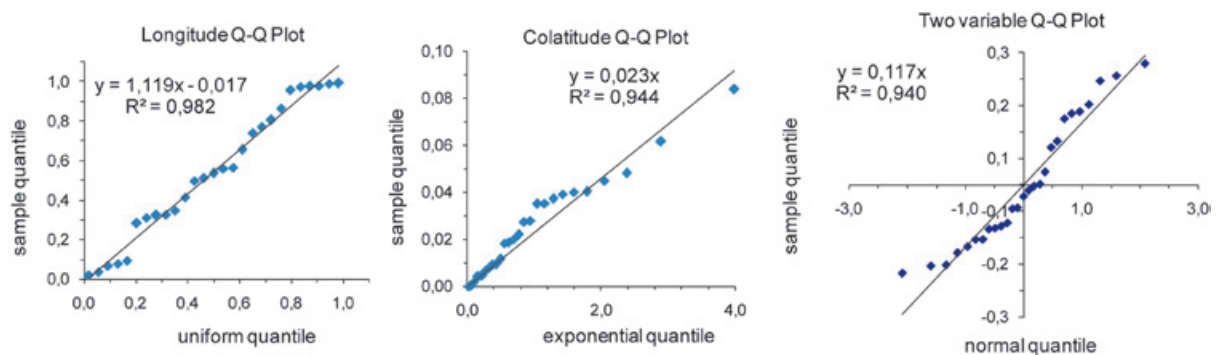


Figure 7. Probability plots for checking goodness of fit of the VGPs to the Fisher distribution.

Table 2. Selected Cretaceous paleomagnetic poles. P= number assigned to paleomagnetic pole (PP) in Figures X; long E, lat S, A95= longitude, latitude and semi angle of 95% confidence of the PPs.

Table 4. Selected Cretaceous PPs.

	Locality	P	Long (°E)	Lat (°S)	α_{95}	Age (Ma)	References
MIS:	Misiones	1	339.1	89.7	4.2	?	Mena <i>et al.</i> (2011)
SPB:	Southern Paraná Basin	2	106.2	84	1.5	~133	Raposo and Ernesto (1995)
CPB:	Central Paraná Basin	3	64.4	84.1	2.3	~132	Raposo and Ernesto (1995)
NPB:	Northern Paraná Basin	4	71.4	83	2.4	132	Ernesto <i>et al.</i> (1999)
PGR:	Ponta Grossa dikes (Brazil)	5	30.3	82.4	2	131–129	Raposo and Ernesto (1995)
CPMP:	Central Paraná Magmatic Province, Brazil	6	197.9	85.7	2.6	133–132	Alva-Valdivia <i>et al.</i> (2003)
CAP:	Central Alkaline Province, Paraguay	7	62.3	85.4	3.1	130–127	Ernesto <i>et al.</i> (1996)
Cba:	Sierras Pampeanas, Córdoba, Argentina	8	75.9	86	3.3	~130	Geuna and Vizán (1998)
FLD:	Florianopolis dikes, Southern Brazil	9	3.3	89.1	2.7	128–119	Raposo <i>et al.</i> (1998)
Baq:	Baqueró Group	10	42.7	88.2	5.5	119	Somoza <i>et al.</i> (2005)
LAd:	Fm. Los Adobes	11	159	87	3.8	130–112	Geuna <i>et al.</i> (2000)
SAEC:	Mean South America Early Cretaceous	12	76.4	84.7	2	135–130	Somoza and Zaffarana (2008)
SAMC1:	Mean South America Middle Cretaceous	13	33.8	89.1	2.4	125–100	Somoza and Zaffarana (2008)
SAMC2:	Mean South America Middle Cretaceous	14	354.1	88.7	2.3	125–100	Somoza and Zaffarana (2008)
SALC:	Mean South America Late Cretaceous	15	345.1	80.6	4.3	85	Somoza and Zaffarana (2008)
UY:	Fm. Arapey	16	95.8	84.8	4.2	~130	Solano <i>et al.</i> (2010)
PY1:	Fm. Alto Paraguay 18	17	359.2	86.2	4.3	~134	Goguitchaichvili <i>et al.</i> , 2012
PY2:	Fm. Alto Paraguay 16	18	34.1	86.1	3.6	~134	Goguitchaichvili <i>et al.</i> , 2013
BC	Reference poles at 135 Ma	19	246.5	83.2	5.5	135	Besse and Courtillot (2002)
PG2	Ponta Grossa	20	222.0	88.1	5		This study

axis for the interval 125 to 100 Ma. Somoza and Zaffrana (2008)'s study also shows that the corresponding paleolatitudes are inconsistent with true polar wander, supporting that the angular discrepancies between the paleomagnetic and fixed hotspot reference frames are due to relative motion of the Atlantic hotspots. The apparent polar standstill of South America has been documented for Eurasia and possibly for the Pacific plate during mid-Cretaceous times (Sager, 2006).

The location of our paleomagnetic poles significantly differs from early Ponta Grossa poles obtained by Raposo and Ernesto (1995). However, these new poles agree well (within uncertainties) with the reference poles given by Besse and Courtillot (2002). This attests that our results should be considered as most precise determination of Cretaceous paleomagnetic pole positions for stable South America.

The VGP's scatter was calculated according to , (Cox, 1969) where S_p is the total angular

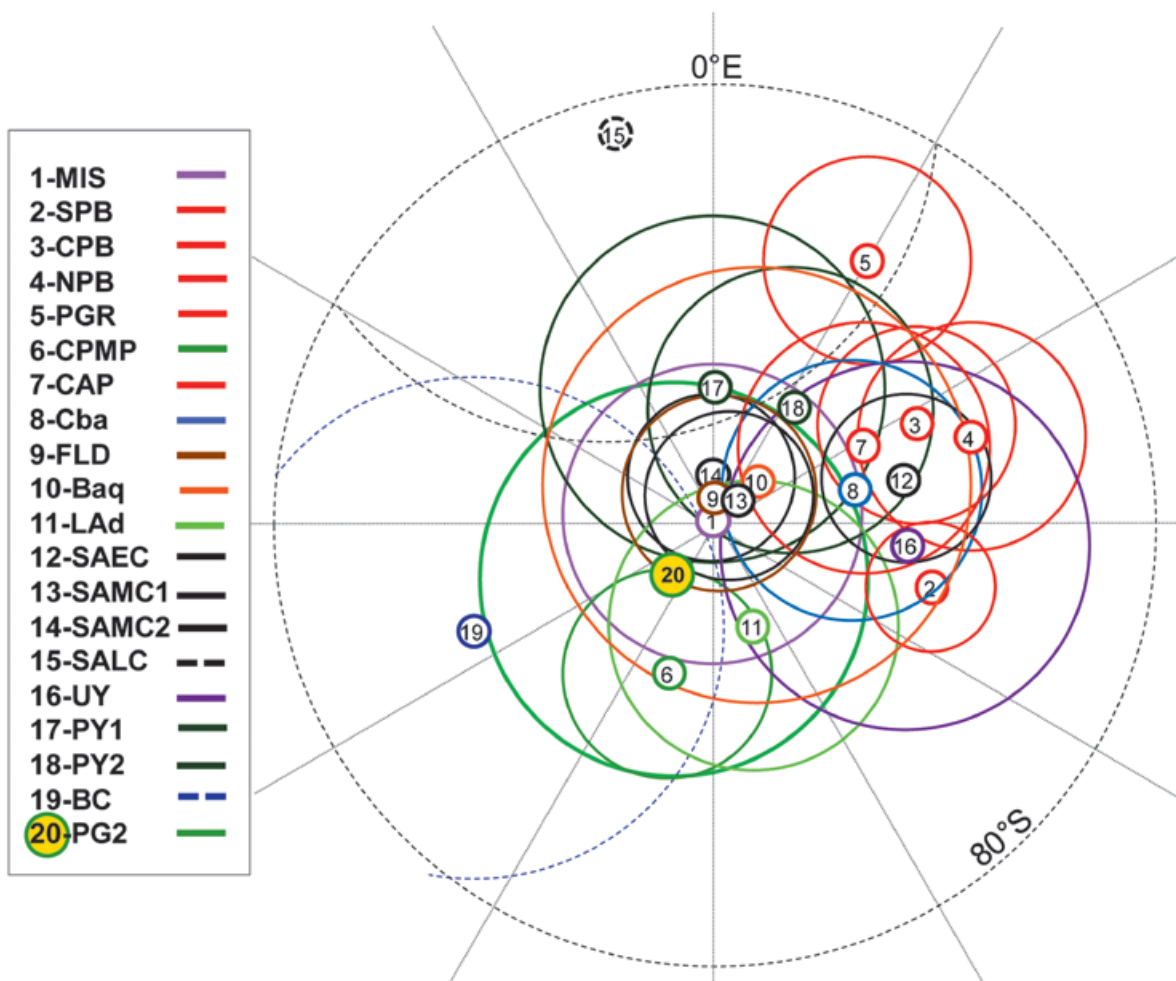


Figure 8. Paleomagnetic poles with 95% confidence circles for Ponta Grossa dikes and previously published Cretaceous PPs with their confidence limits. 1. Misiones (Mena *et al.*, 2011), 2. Southern Paraná Basin (Raposo and Ernesto, 1995), 3. Central Paraná Basin (Raposo and Ernesto, 1995), 4. Northern Paraná Basin (Ernesto *et al.*, 1999), 5. Ponta Grossa Dikes (Raposo and Ernesto, 1995), 6. Central Paraná Magmatic Province (Alva-Valdivia *et al.*, 2003), 7. Central Alkaline Province, Paraguay (Ernesto *et al.*, 1996), 8. Sierras Pampeanas (Geuna and Vizán, 1998), 9. Florianópolis Dikes (Raposo *et al.*, 1998), 10. Baqueró Group (Somoza *et al.*, 2005), 11. Los Adobos (Geuna *et al.*, 2000), 12. Mean Early Cretaceous (Somoza and Zaffrana, 2008), 13. Mean Middle Cretaceous (Somoza and Zaffrana, 2008), 14. Mean Cretaceous (Somoza and Zaffrana, 2008), 15. Mean Late Cretaceous, (Somoza and Zaffrana, 2008). 16. Araucary (Solano *et al.*, 2010), 17. Alto Paraguay 18 (Goguitchaichvili *et al.*, 2012), 18. Alto Paraguay 16 (Goguitchaichvili *et al.*, 2013), 19. Reference Poles at 135 Ma (Besse and Courtillot, 2002), Ponta Grossa (This study).

dispersion, D_i is the angular distance of the i -th VGP to the PP and N is the number of sites used in the calculation. Thereby $S_p=12.9^\circ$. The paleosecular variation (PSV) was estimate according to McElhinny and McFadden (1997) by where $S_w = 81/K^{1/2}$ is the within-site scatter, K being Fisher concentration parameter for poles estimate from direction concentration parameter (k) according to $K = 8k (5 + 18\sin^2 \lambda + 9\sin^4 \lambda)$, λ is the site paleolatitude, and n_i is the number of directions from i -th site. The calculated PSV is $S_b=12.3^\circ$, with a 95% confidence interval $[10.3^\circ; 15.2^\circ]$, which agrees well with the expected value for this latitude. The S_b values here obtained agree well with the selected data reported for the Cretaceous Normal Superchron (Biggin *et al.*, 2008) between 125 to 84 Ma. This supports that the Earth's magnetic field maintained similar characteristics before and during CNS, in disagreement with previous studies.

Acknowledgment

The financial support for this work was provided by UNAM DGAPA PAPIIT projects IN-105215, 107114 and IA-104215, as well by CONACYT project 129653. This work was also funded by CGL2012-32149 project (Ministerio de Economía y Competitividad, Spain) and the European Regional Development Fund (ERDF).

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